



Mafic explosive volcanic activity- controlling factors and conditions

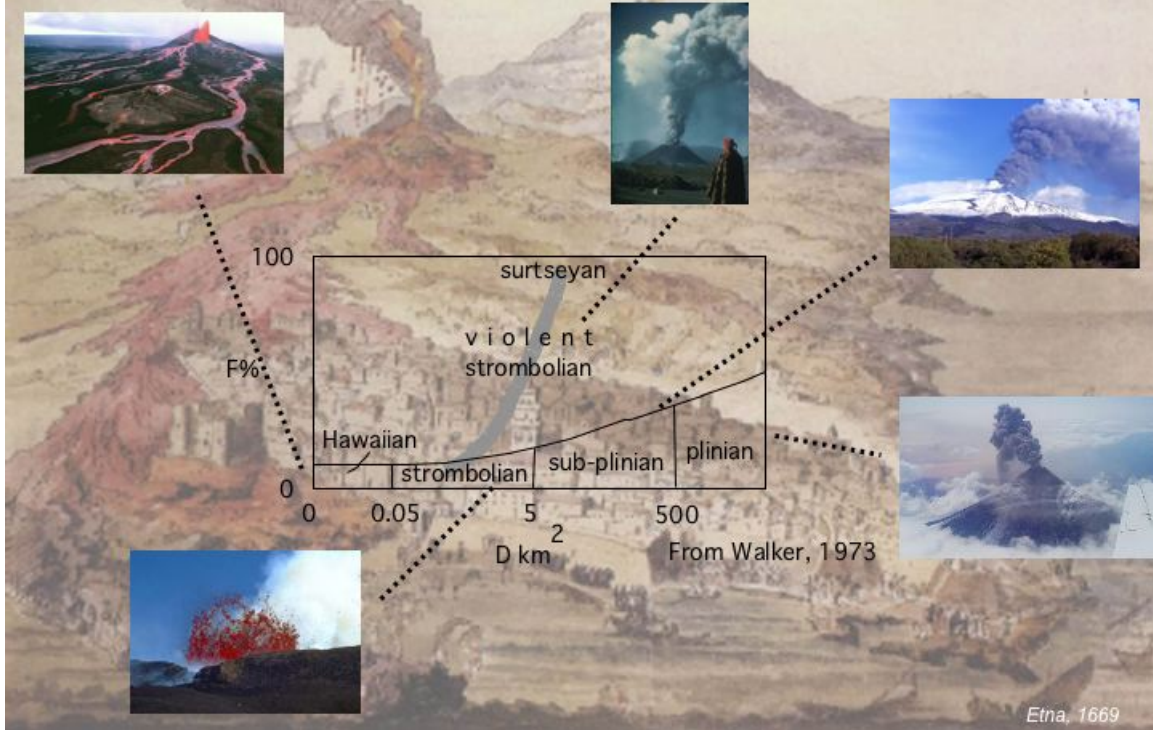
L. Pioli, Université de Genève

Dr. Atl, El Parícutin

Misconceptions on basaltic eruptions

- Basaltic magmas are typically volatile-poor
- Basaltic eruptions rarely produce voluminous amounts of ash and lapilli
- 'Explosive' basaltic eruptions are rare events

An empirical classification



Length, frequency, characteristics

- A few hours to several years
- Single explosions spaced up to $< 1s$
- Mild to strong fluctuations in intensity
- Some explosions coexist with lava flows
- Commonly eruptions start from fractures
- Multiple vents



Hawaiian eruptions



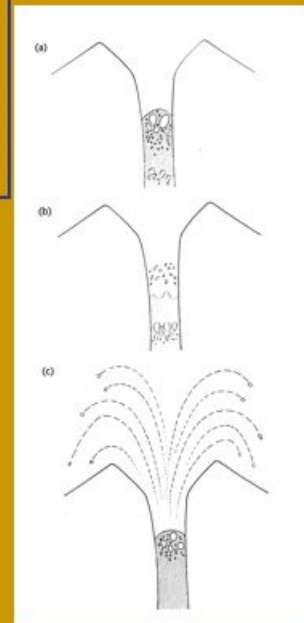
From fissures to central conduit
Fire fountains or curtains
 $H = 100-500$ (1600) m
 $MER = 50-1000 \text{ m}^3/\text{s}$
Gas/magma $\sim 70:1$
Associated to lava flows
Mainly cone-forming

Puu Oo, Hawaii

Strombolian eruptions



$H = 100-500$ m
Fountain structure
 $MER = 10^{-3} - 10 \text{ m}^3/\text{s}$
Gas/magma $\sim 10000:1$
Explosion frequency:
 $10^0 - 10^2/\text{hour}$
Typically cone forming

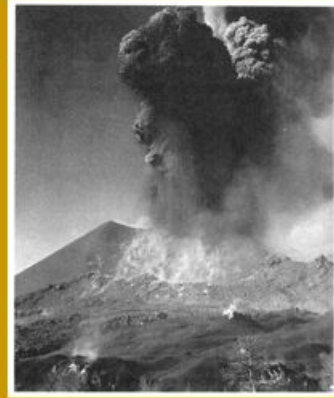


Increasing explosivity...

- Violent strombolian
- Surtseyan
- Subplinian
- Plinian



Etna (2002-2005)



Paricutin-Mexico (1943)

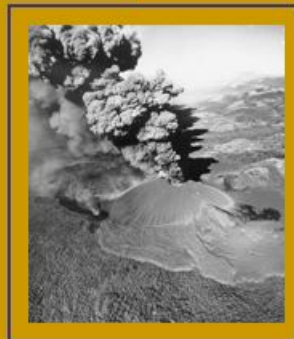


Violent strombolian activity

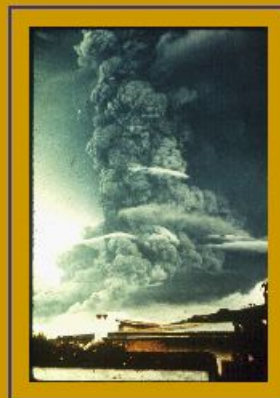


Paricutin,
Mexico
(1943-1951)

- Both cone and sheet-forming
- Associated with lava flows
- Pulsatory, non sustained dynamics
- High ash-charged columns



Cerro Negro, Nicaragua (1968)



Fuego,
Guatemala
(1974)

Violent strombolian activity



- Lava flow emission typically contemporaneous with explosions, usually from lateral vents
- Shallow gas segregation

Tolbachick, Kamchatka (1975)

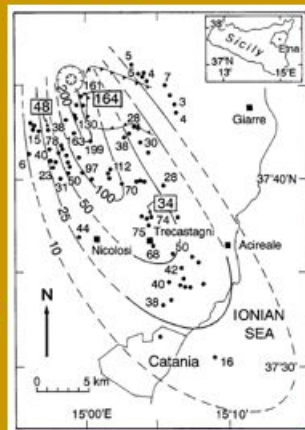
Plinian-subplinian activity

Eruption column height: 20-32 km (Plinian)
>8-10km (Subplinian)



Shishaldin,
Alaska (1999)

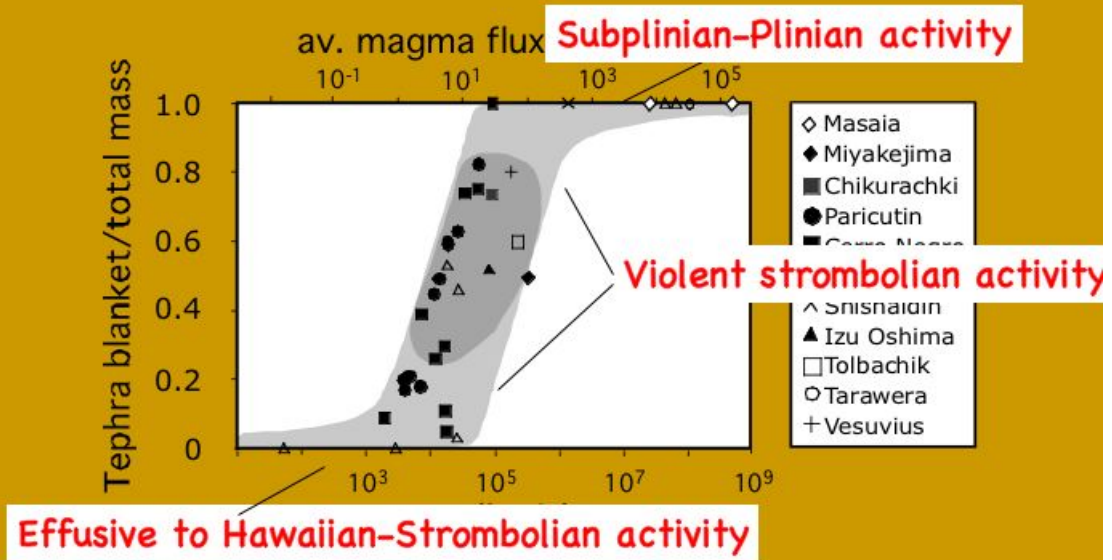
Etna (122 bc)



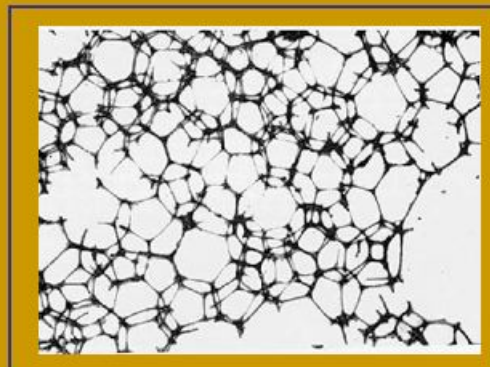
KNOWN PLINIAN ERUPTIONS:

Etna, 122 b.c.; Tarawera, 1886; Masaya, 60 and 2 ka b.p.

Mild to high explosivity transition



Products of eruptions-Hawaiian



- reticulite

Products-Hawaiian to Strombolian



spatter



Products -Hawaiian to Strombolian



bombs





Strombolian scoria

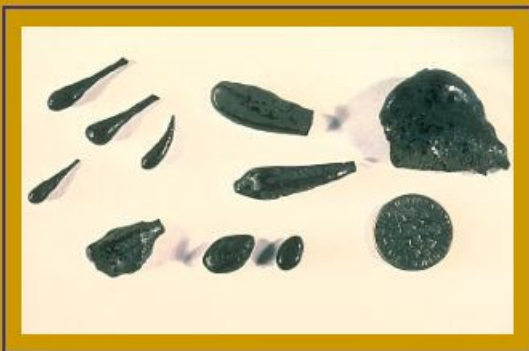


Strombolian activity- scoria cones



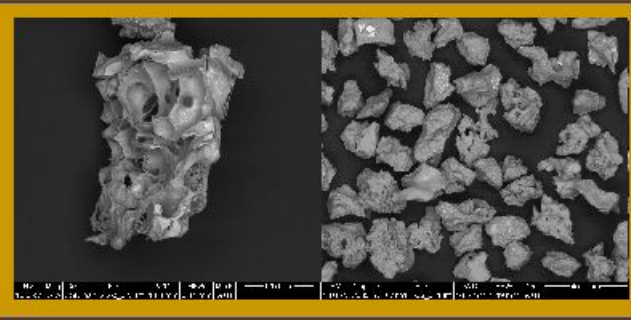
- Scoria and spatter accumulate around the vent
- Grain flow

Fine particles -Hawaiian

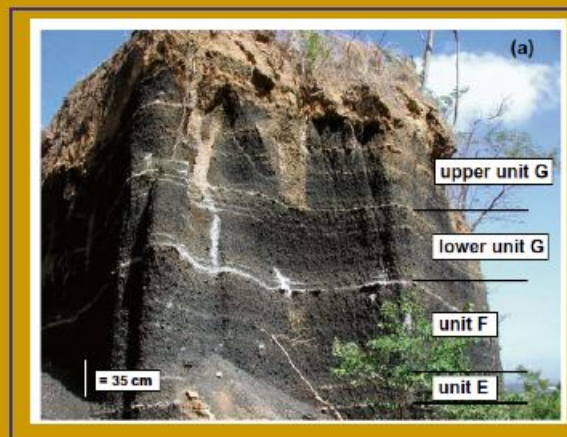
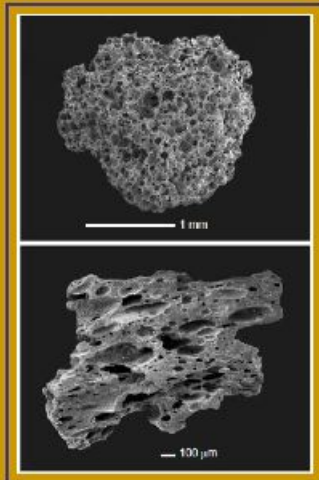


- Pele's Hair
- Pele's tears

Ash - higher explosivity



Plinian activity- tephra blanket

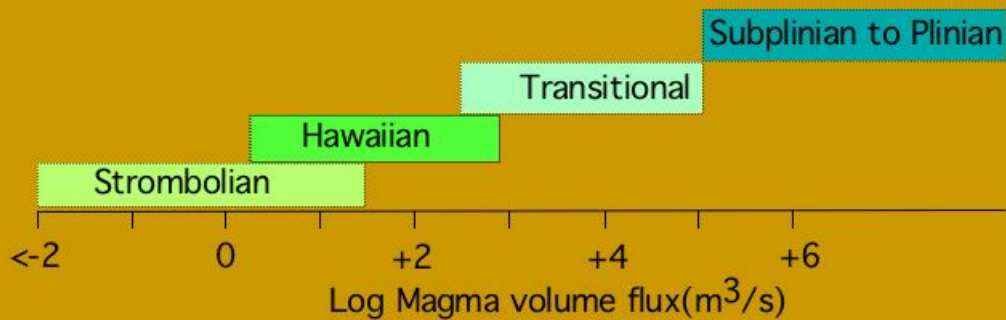


Fontana Lapilli,
Masaya volcano
(Nicaragua)

Main factors controlling explosivity of mafic magmas

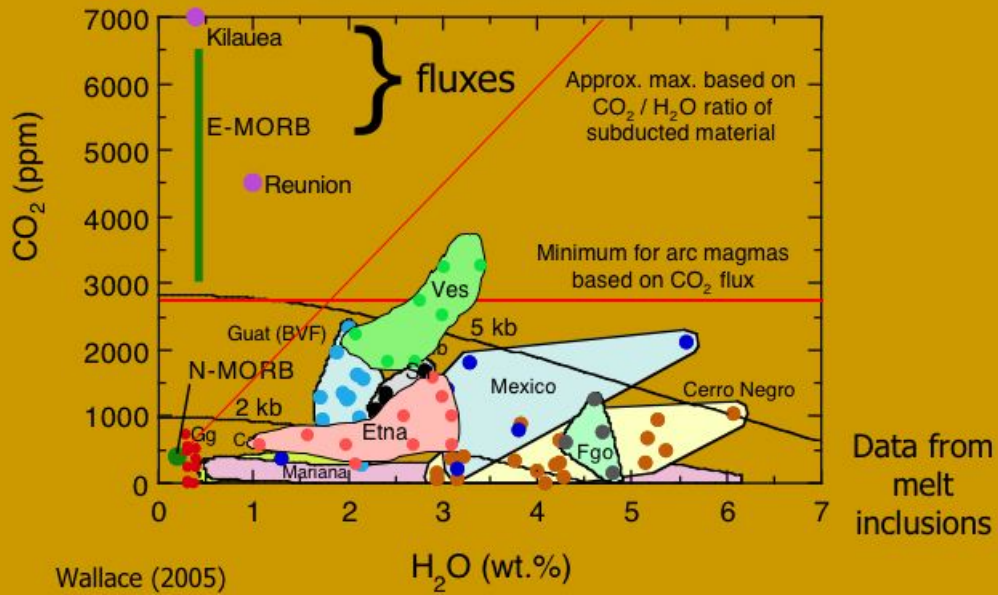
- Magma rise rate/magma supply rate
- Volatile content; crystal content
- Conduit/dike geometry

Mass eruption rate and eruptive style

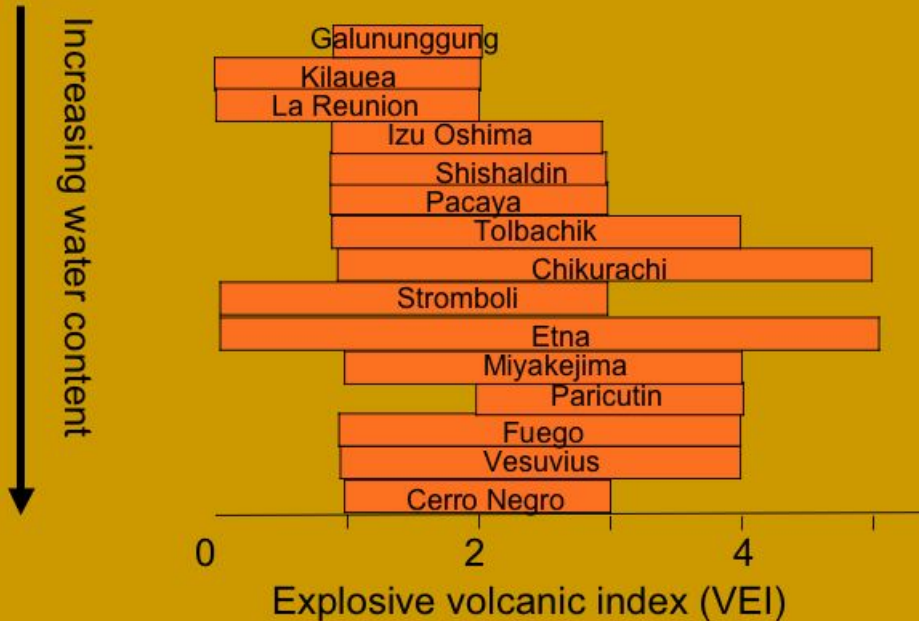


Main volatiles content

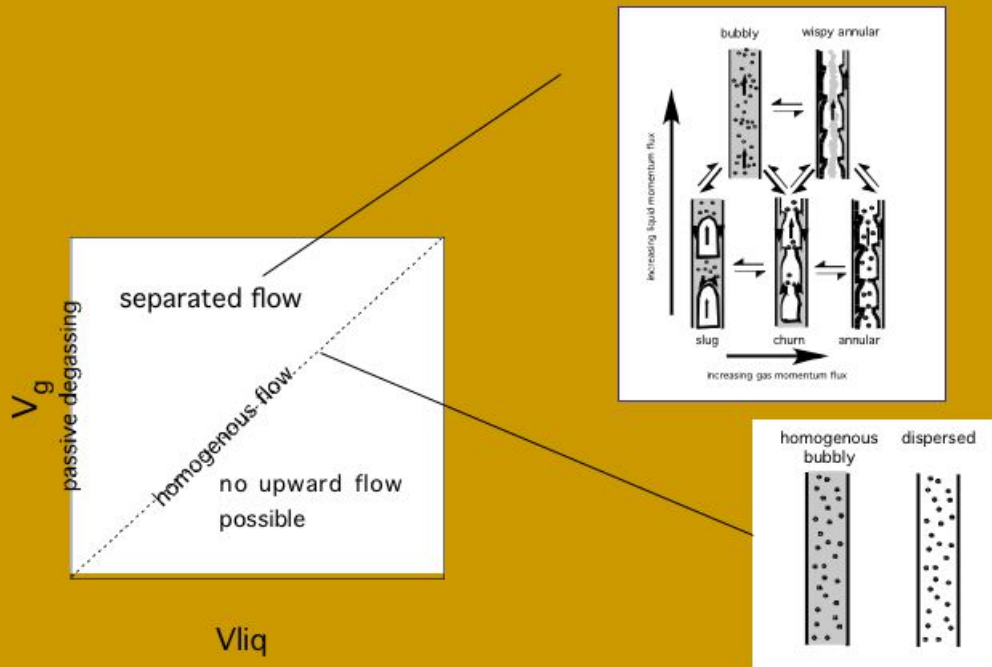
H₂O and CO₂ in Basaltic Magmas



Basalt explosivity and magma volatile content



Two-phase flow



Bubble rise within the melt

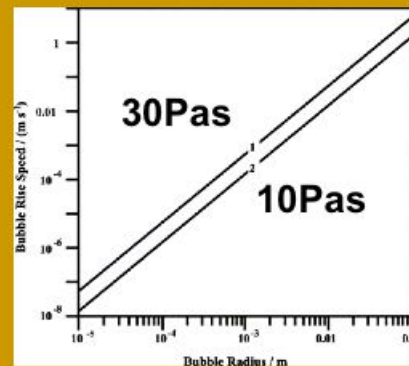
Basalts have viscosities ranging from 10 - 10^3 Pa*s depending on composition

- At small Re single bubbles rise obeying the law:

$$U_{\infty} = k \frac{\rho g d^2}{\eta}$$

- With $2/9$ (Hadamard -Rybczynski law) or $k=1/3$ (Stokes law) with increasing Re
- Large bubbles rise independently of their size

$$v = 1.54 \left[\frac{\sigma g}{\rho_l} \right]^{0.25}$$

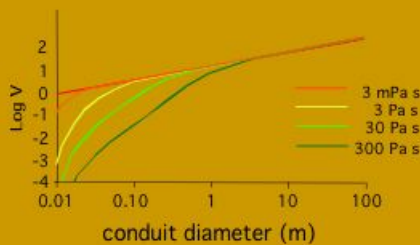
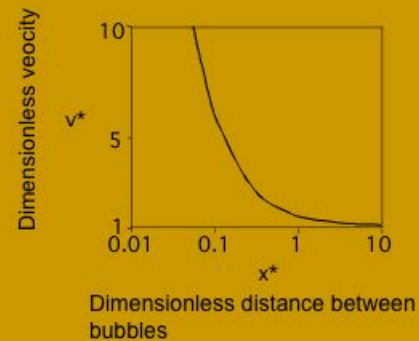


Wilson and Head (1981)

- But.. Single non interacting bubbles rarely exist in nature!

Bubble rise within the melt

In a train of bubbles, single bubble rise is affected by the wake of the previous bubble



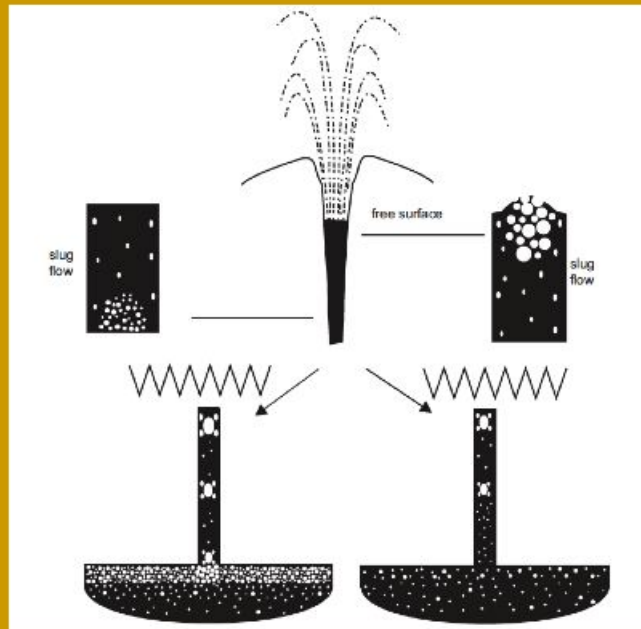
Bubbles almost as large as the conduit-

Taylor bubbles rise independently of their diameter and magma viscosity above $d=1$ m

Implication of separate flow for eruptive activity

- Gas segregation
- Effects on explosivity
- Overpressure buildup and bubble burst

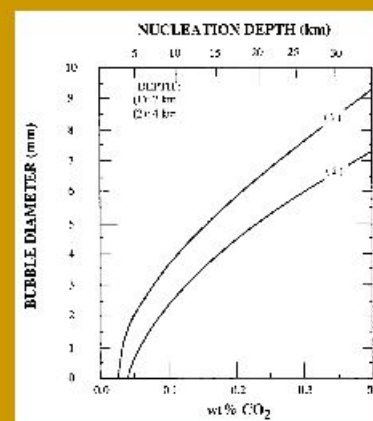
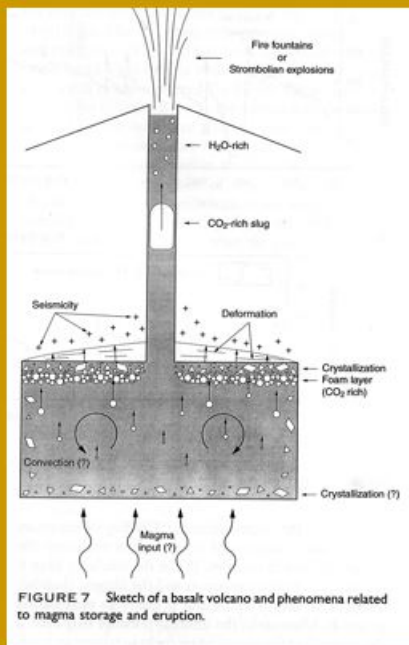
Two models for separate flow...



Houghton and Gonnermann (2009)

Foam collapse

(Vergnolle and Jaupart, 1986)



Vergnolle and Jaupart (1990)

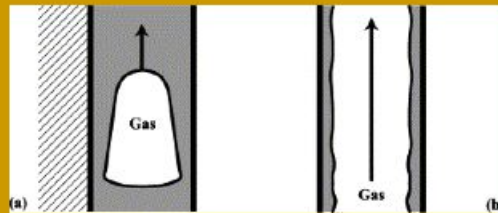
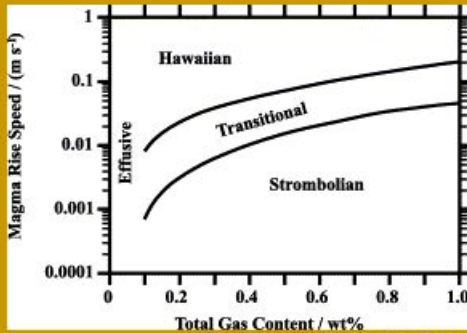
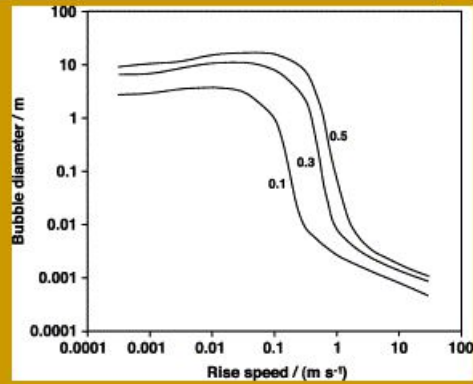
Eruptions are CO₂-driven

Vergnolle and Mangan (2000)

Magma rise speed model

(Head and Wilson, 1994)

- Magma rise speed dependent dynamics
- Threshold $V=0.1$ m/s



Parfitt (2004)

Two-phase flow of basaltic magmas in vertical cylindrical conduits

